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OCEANOGRAPHIC INSTRUMENTATION



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1. INTRODUCTION

Oceanographic instruments measure or sample various physical, chemical and biological quantities in the water column. There is often more than one way to measure a particular quantity (i.e. temperature, currents) and there are many quantities to measure hence there are many instruments. This course will attempt to cover the more common instruments used by the oceanographer.

We will first start by listing which quantities commonly measured these include:

*Depth (metres) *Temperature (degree Celsius) *Salinity (practical salinity units of chlorine, sodium, sulfate, magnesium and potassium) *Oxygen, *Phosphate, *Nitrate, *Silicate, *pH *Density (kilos per cubic meter) *Water Clarity (Forel scale to measure colour) *Sound or Ambient Noise (hertz) *Sound Speed *Bioluminescence *Seabed sediment *Current *Waves

Some physical properties of seawater are conservative (i.e. temperature and salinity), this means that away from the surface the only way they can change is by mixing.

All of the above quantities are normally accompanied by a measurement of date, time and position i.e latitude and longitude. Measurement of time and position are made with the Global Positioning System (GPS) and are relative to the GMT time. Accuracy is of the oder of +/- 10-15m.

What follows is description of the methods and instruments commonly used to measure each of the above quantities.

2. QUANTITIES MEASURED AND INSTRUMENTS USED

2.1 DEPTH

The measurement of the depth regards the instrument or the sea-floor and is fundamental to reference the location of the measurement.

The simplest way to measure the depth of the measurement is to mark the wire or rope to which the instrument is attached and measure the length used.

The more accurate way is to measure the water pressure at the instrument level and convert this to a depth.

The simplest way to measure sea-bed depth is to lower a weight with a wire or rope and measure the length used till the weight has reached the bottom.

Yet another is to use an *ECHO SOUNDER* and time how long takes for the sound pulse to echo back. Two instruments are the **SONAR** (sound navigation and ranging) and the **SOFAR** (sound fixing and ranging).



Figure 1. Velocimeter, sonic frequency system

A *CONDUCTIVITY, TEMPERATURE AND DEPTH (CTD)* instrument uses a strain gauge pressure transducer to sense depth.

2.2 TEMPERATURE

There are several ways of measuring temperature, the simplest being a mercury-in-glass or *STEM THERMOMETER*. A stem thermometer is commonly used to measure sea surface temperature by placing it in a bucket of sea water.

A **PROTECTED REVERSING THERMOMETER** is special kind of stem thermometer which measures sub-surface temperature [Figure 1]. 'Protected' means that the thermometer is isolated from water pressure. The reversing thermometer is attached to a **WATER SAMPLING BOTTLE**. When the sampling bottle is closed the thermometer is inverted and, as a result of its construction the mercury 'breaks' at a particular point and runs down the other end of the capillary to record the temperature at the depth of the reversal. After the thermometer has been reversed it becomes almost insensitive to subsequent changes of temperature and it is read when it is bought back on deck. After corrections for scale errors and for the small change in reading due to any difference between the insitu temperature and that on the deck, the reversing thermometer is accurate to about +/- 0.02 C in routine use.



Figure 2. Nansen reversing water bottle

An **UNPROTECTED REVERSING THERMOMETER** is as described above except that it is exposed to water pressure. The water pressure compresses the glass in the bulb and causes the thermometer to indicate a higher temperature than the protected thermometer. The difference between the two thermometers is a measure of the compression of the glass which depends on a known compressibility and upon pressure i.e the *depth*.

Therefore a protected/unprotected reversing thermometer pair is used to measure both temperature and depth. The depth is then used to reference the depth at which the water bottle sample was taken. Depth measured in this way is accurate to +/- 0.5% of depth or +/- 5m whichever is greater.



Figure 3. Reversing thermometer

Another widely used temperature sensor is the *THERMISTOR*. The thermistor is usually a small piece of electrically conductive material such as a resistor. The thermistor relies upon measuring electrical resistance which is directly or inversely proportional to temperature.

Thermistors are used on **CTD** (conductivity, temperature and depth) and **EXPENDABLE BATHYTHERMOGRAPH** (XBT) instruments.



Figure 4. CTD (conductivity, temperature and depth)

The CTD instrument is composed by a Rosette with weights, 12 to 24 Nisking bottles, reversing protected and unprotected thermometers and other instruments such as a fluorometer

and a transmissometer. The water samples are used for collection and used for calibration and chemical/nutrient analysis. The instrument provides real-time sections or profiles of temperature, salinity, density, geostrophic currents and dissolved oxygen. The temperature/salinity diagrams are used for water mass identification.





Figure 5. XBT (Expendable bathythermograph) launching

The XBT is composed by a 'torpedo'-shaped probe which contains a thermistor and a very fine coiled copper wire which is unreeled both from the probe and from the canister from which it is deployed. It is launched from the ship being either hand-held, deck-mounted or hull-mounted. The depth is calculated from the elapsed time and the expected fall rate (~6.5 m/s) and comes into 4 types (T4 ~460m, T7 ~760 m, T10 ~200m, TDeep ~760). This instrument provides actually a large body of data (global dataset) over many seasons/years for statistical evaluation of climatic changes.



Figure 6. XBT probe

THERMISTOR 'CHAINS' consisting of a cable with a number of thermistor elements at intervals are sometimes moored along with current meters to record temperature at a number of depths. A 'data logger' samples each thermistor at regular intervals and records temperatures as a function of time.

A **SEASOR** is an 'small aeroplane' looking instrument that is lowered in the water and moored by the ship at a known depth or at the surface, with the aid of a conducting cable the data is real-time monitored and produces typical sections such as virtual CTD profiles. (see picture, slide from jane)



Figure 7. Seasoar

Sea surface temperature may be inferred by measuring the infra-red radiation emitted by the sea surface. This method is used by airborne platforms such as *SATELLITES*. Use of satellites or other 'remote-sensing' methods permits rapid measurement over a wide area.

2.3 SALINITY

Salinity is a measure of the quantity of salt in a volume of sea water. It has the standard units of Practical Salinity Units (psu). In very approximate terms a salinity of 35 psu means there are 35 grams of salt in a kilogram of water.

Salinity is an important quantity to measure as salinity in conjunction with temperature and depth/pressure enables calculation of the density of the sea water and also sound speed.

Salinity in combination with temperature also allows the oceanographer to label particular water masses in the ocean and study the movement of these water masses and hence infer ocean currents.

Salinity is obtained by first measuring the electrical conductivity of the sea water at a known temperature and pressure. Using an internationally standard formula the conductivity is then converted to a salinity.

Salinity may be measured from a water bottle sample on the ship using a laboratory *SALINOMETER* or in-situ using a conductivity sensor such as on a **CTD**.

The salinometer measures a conductivity ratio using a conductance cell and conductance bridge which is balanced manually. The ratio measured is between a standard sample of sea water at 35.000 psu and the sample. Salinity is then obtained from the conductivity ratio and temperature using standard tables or formulas.

CTD's commonly use a 4 electrode cell through which water flows as the cell is moved through the water column. The cell measures conductivity which is then converted to salinity. Figure 3 shows a CTD instrument with a array of water sampling bottles and reversing thermometer attached.

2.4 OXYGEN, PHOSPHATE, SILICATE, NITRATE, pH

Oxygen content of sea water is commonly measured in conjunction with CTD measurements or water bottle sampling. CTD units sometimes have a oxygen sensor attached enabling an insitu measurement of oxygen, but these measurements produce noisy results which need to be filtered.

Oxygen content from water samples is determined by chemical titration. The titration method is often more reliable than using an in-situ sensor which may suffer from drift or calibration problems. Figure 8 shows water samples being taken from an array of water bottles attached to a CTD unit.



Figure 8. Wet Laboratory of RV Franklin, Water Bottle Sampling. Oxygen bottle being filled at right, Salinity bottle at left.

Phosphate, silicate and nitrate are determined by chemical means from water samples. Other chemical parameters such as pH may be determined from water samples.

2.5 WATER CLARITY

Water clarity may be taken as a measure of the degree of transmission of visible light through the sea so are taken whenever possible during daylight hours. The more 'turbid' the water the less light transmitted. Water clarity has allways been measured to estimate the silt run-off from rivers, monitoring pollution streams ans monitoring algae growth.

The simplest device for measuring water clarity is the *SECCHI DISC* which consists of a 30 cm in diameter white plate which is lowered into the water (a lead weight is suspended under the disc to ensure that it will sink rapidly and vertically) and depth at which it lost to sight noted. The greater the depth at which the disc is no longer visible the clearer the water. During the procedure the colour of the disc is classified with the Florel-Ule scale. The secchi disc is only a semi-quantitative device but being simple and a low-cost device it is often used.



Figure 9. Secchi Disc

A more quantitative device is the *TRANSMISSOMETER* which measures the attenuation of the beam of light of known wavelength over a fixed path length. The transmissometer uses a light source and a photo-electric-cell and measures the beam attenuation co-efficient 'C' over a direct path from the source to the photo-cell. The beam attenuation co-efficient is a function of the shapes and amount of particulate material in the water.

Another instrument is the *NEPHELOMETER* which measures light scattered through an angle rather than from a direct path from the light source to the photocell. Nephelometer measurements have a more direct relationship with the amount of suspended solids in the water than a transmissometer and is thus used when one wants determine sediment concentrations in the water i.e in grams/volume units, The nephelometer is useful in muddy highly turbid water such as near coastal estuaries, rivers and near sea bottom (nepheloid layer).

2.6 SOUND

Sound is measured in the sea to support a number of applications most notably for Naval sonar operations, geophysical studies and also for studying sounds emitted by marine life.

The basic instrument used to measure sound is the *HYDROPHONE*. A hydrophone is a transducer which converts sound energy (pressure) into electrical energy (current). Some substances such as quartz or certain ceramics when placed under pressure acquire an electrical charge or voltage across the crystal surface. This behaviour is called 'piezo-electricity'. Thus the hydrophone consists of sensor made of a crystalline or ceramic substance.

The basic unit of sound measurement is the 'decibel' which is measure of the pressure exerted by the sound wave or measure of the sound intensity. The other important quantity is the frequency of the wave measured in 'Hertz'.

Sound in the sea may be thought of consisting of many super-imposed waves at varying levels of intensity and frequency resulting in a complex wave arriving at the sensor. Using mathematical techniques the complex wave is decomposed into waves of discrete frequency and intensity. The result of this analysis is a graph of intensity (decibels) versus frequency (Hertz) called a 'spectrum'.

The sound or background noise is frequently referred to as 'ambient noise'.

Many hydrophones consist of more than one sensor element or an 'array'. The advantages of an array over a single hydrophone are several. First the array is more sensitive, since a number of elements will generate more voltage (if connected in series) or more current (if connected in parallel) than a single element exposed to the same sound field. Second, the array possesses directional properties that enable it to discriminate from sounds arriving from different directions. Third, the array has an improved signal-to-noise ratio since it discriminates against isotropic noise in favour of a signal arriving in the direction that an element of the array is pointing.

Sonar domes fixed to the hull of naval vessels have a cylindrical array. A recent development is the 'towed array' which of a flexible line of hydrophone elements towed from a ship.

Another type of hydrophone is used on a *SONOBUOY* which is dropped from an airborne platform to record ambient noise and locate ships or submarines. A sonobuoy is compact expendable device containing a small radio transmitter for relaying signals picked up by the hydrophone.

2.7 SOUND SPEED

Sound speed is a measure of how quickly sound propagates through the ocean. Sound speed is needed for measuring the water depth with *ECHO SOUNDERS* since the water depth is calculated by multiplying the mean sound speed in the water column by half the time that it takes for the sound to echo back from the sea bed.

The basic means of measuring sound speed is to measure the time it takes for a pulse of sound to travel a known distance. The sensor consists a sound source a metal reflector and receiver.

The sensor may be attached to a probe which is lowered into the water or on an expendable instrument such as the *EXPENDABLE SOUND VELOCITY (XSV)* probe. The XSV works on simple principle to the XBT.

Sound speed may be calculated from a well known relation between sound speed, temperature, salinity and depth. So in the absence of direct sound speed measurements sound speed can be estimated quite accurately provided temperature, salinity and depth are known.

2.8 CURRENTS

Measurement of ocean current is fundamental to a general understanding the ocean. Ocean currents transport heat and effect climate and weather. Currents transport marine life and sediments. Currents also effect the passage of ships and also the oceanographic instruments themselves.

Current is a vector quantity having both a direction and a speed. Therefore a current instrument must measure both direction and speed.

There are different basic ways of measuring a current, these are:

- 1. The oldest method of tracing a current is by the use of a *DRIFT BOTTLE* or *CARD*. These floating items are designed so that their movement is determined only by surface currents and contain a request to the finder to notify the time and place of recovery
- 2. A *MOORED CURRENT METER* in the sea produces measurements of speed and direction at fixed time intervals.
- 3. Attach the current meter to a moving platform such as a ship and measure the relative current. The absolute current is then determined by such vector subtracting the motion of the ships over the ground determined by navigation. This method is used with the *ACOUSTIC DOPPLER CURRENT PROFILER (ADCP)*. An ADCP may also be moored at a fixed location on the sea-bed.
- 4. Let an object flow or float with the current and observe its motion i.e the change of position at fixed time intervals. For example a *DRIFTING BUOY*.
- 5. A *DYNAMIC METHOD* recognises that the surface of the sea is not level but has hills and valleys. Therefore countour maps may be drawn of its surface showing these elevations and depressions. Taking into account that in the southern hemisphere ocean currents flow in an anti-clockwise direction around areas of high sea level, a map of sea flow of an area can be drawn from a knowledge of the topography of the surface. These contours are not visible by the naked eye but are correlated to the density of the water, the water below a hill being of lower density than that below a valley. The density of the water instead of being directly measured is computed from the temperature and the salinity of the ocean water.

Owing to turbulence current is a quantity which often fluctuates over short periods of time and usually averaging or smoothing is required to present the data in useful form. This process of averaging is termed 'vector averaging'. Current meters which use vector averaging sample the current at short intervals, say once every 2 seconds, and then compute the average over a longer interval, say 10 minutes. The resulting output of the meter is a time series of current at the averaging interval.

There are 4 common techniques of current measurement as follow

2.8.1 Propellor Type Meters

These instruments rely a measuring the counting (metering) the rotation of a rotor or propeller placed in the current. Direction is sensed by a vane which is orientates parallel to the current The orientation of the vane is then referenced against a magnetic compass. An example of such a current meter is the 'Aanderaa' model.



Figure 10. Standard type of current meter and Ekman current meter

2.8.2 Electromagnetic Current Meters

These instruments use the principle that a voltage will be induced by a conductor which moves across a magnetic field. The conductor in this case is sea-water which readily conducts a current due to salt ions.

Electromagnetic current meters consist of a pairs of electrodes, an internally generated magnetic field and a flux gate compass. Water flows through the magnetic field thereby producing a voltage which is proportional to the current speed. The current amplitude is measured along the axis of each electrode as well as the compass output. Relative current direction, corrected with compass data and current amplitude are computed internally. The results are then stored in two absolute current vector components (North and East). The components are then vector averaged over a user programmed averaging interval.

2.8.3 Acoustic Current Meters

Acoustic current meters utilise the Doppler effect named after the Austrian physicist C. J. Doppler. The Doppler effect is the change in frequency of sound reflected by a stationary object relative to the frequency at which it is moving.

The **ACOUSTIC DOPPLER CURRENT PROFILER** (**ADCP**) emits a beam of sound of known frequency which reflects off small particles moving with the water. The beam reflected back to the receiver will have a change in frequency proportional to the speed of the particles and thus the current speed.

One sound beam will give the component of current in the direction of the beam. However three orthogonal components are needed to get the true current vector so the ADCP utilises more than one beam. Four beams are typically used to obtain a redundant velocity measurement for data checking and improved instrument reliability.

ADCP's measure water speed at multiple water depths or 'range cells' along the path of the acoustic beams. This is achieved by periodically transmitting short pulses (typically 1-50 milliseconds) of sound, then making multiple measurements of the frequency of the echoes at discrete time intervals after the initial sound pulse.

In this way a depth profile of the current is assembled after some averaging processes. On a moving platform such as a ship the motion of the ship must be subtracted from the ADCP relative current to get the true current. Usually the motion of ship is determined by using GPS navigation which offers the best means of accurately determining the ships motion at frequent intervals.

The end result of the ADCP data processing is a time series of current versus depth profile along the track of ship. In the case of moored ADCP it is time series at a fixed point.

2.8.4 Drifting Buoys

A drifting buoy is a device which floats on the ocean surface to follow the current. The buoy usually includes a 'drouge' which is device such as a parachute or sheet to drag the buoy with current. The drouge is attached to the buoy below the water.

Some method is needed of tracking the buoys position. Position tracking may be done using the following methods:

- 1. Visual sighting with a bearing compass.
- 2. Attaching a metal target/reflector to the buoy so it can traced by radar.
- 3. Using a radio transmitter and tracking the buoy location with a radio receiver such as satellite.

Drifting buoys may have other sensors attached such as thermistors which measure sea surface temperature, a pressure gauge to measure air pressure and an anemometer which measures wind speed and direction.

Another type of lagrangian device is a spherical free-falling acoustic device with a 'skirt' which reduces turbulence as it falls. This device with the help of a triangular configuration of

transponders on the seabed can be located with high precision while sinking and drifting with the currents. This device produces a profile with 'real' current components with depth and after each profile, so after it has reached the seabed, it reached the surface again awaiting for collection.



Figure 11. Pegasus velocity recorder

Other means of tracing 3-D movements of water are acquired with the aid of tracers such as dye and pollutants.

2.9 WAVES AND TIDES

The most common instrument used for measuring waves is the *WAVERIDER BUOY* (figure 12). A waverider buoy is a buoy which follows the movements of the water surface.



Figure 12. Waverider Buoy

The wave height is measured with an *ACCELEROMETER* which measures the vertical acceleration of the water. The acceleration is integrated to produce a time series of displacement of the sea surface i.e a time series of wave height. The buoy is typically attached to a flexible mooring line which in turn is attached to heavy weight on the sea bed.

The wave height signal is usually transmitted by radio signal to a ship or shore station from an antenna attached to the buoy.

Tides are most commonly measured with a TIDE GAUGE.

2.10 SEABED SAMPLING

Seabed information is very usefull to estimate anchorage safety locations, offshore engineering (such as siting of platforms, beacons and sea walls), mineral exploration, fishing, minewarfare, submarine operations and sonar acoustic performance. The classification of the bottom type relies on a geological/scientific classification, and hydrographic classification (description of the most predominant components, the grain size grading), the nature of the sea floor (washed materials, erosion, biological sediments).



Figure 13.

Seabed sampling can be obtained by different means. Leadlines, Grabs, Snappers and Scoops, Corers, Dredges, Divers, and Remotedly Operated Vehicles and Submersibles.

LEADLINES are weights armed with a sticky substance to which the particles adhere and to which heavier objects leave an impression. The advantages of the leadline is that it is cheap and simple to operate. The disadvantages are that the heavier materials might not be detected, only the very top surface layer is sampled and the latter is disturbed when collected

GRABS, SNAPPERS AND SCOOPS are used for collecting medium size samples from the surface and immediate sub-surface layer of the sea floor. They usually comprise a bucket or scoop which is activated on hitting the sea floor. Some are spring-loaded, others close when they are raised off the seabed. The disadvantages of the grabs are that they are not suitable to collect soft or liquid muds as the sample is often washed out of the bucket before it reaches the surface.



Figure 14. Grab samplers

The **SHIPEK GRAB** consist of two concentric half cilinders that on striking the sea floor rotate through 180 deg. During this rotation the bucket scoops a sample from the seabed and then remains closed whilst the grab is hauled to the surface. This type of grab is most effective on unconsolidated sediments but the impact of the grab on the consolidated and compacted seabed can be liable to make it bounce and to cause only a superficial grab.

CORERS are used to obtain an undisturbed verical sample of the sea floor. They comprise a tube or box shaped cutting mechanism similar to an apple corer or pastry cutter. They are driven into the seafloor and when whithdrawn retain the undisturbed sample of the sediment layers by the means of different methods such as vacuum suction and shutters. Corers can be drawn in the seafloor by means of their own weight, explosives, pneumatics or mechanical vibration. Cores contain liners which allow the sample to be removed and stored with the minimum disturbance.



Figure 15. Corer

DREDGES are designed to drag along the sea floor collecting loose materials and sediments. They often incorporate a filter which allows smaller sediments to pass through. Samples are oviously disturbed but do reflect the seabed materials over a reasonable large area.

DIVERS inspection allow a positive identification of the shallow sea floor. Large as well as small features can be identified.

2.11 BIOLUMINESCENCE

Bioluminescence is the emission of light by living organisms. The process is such that light is radiated but very little heat is emitted. The process is an animal production and approximately 240 groups of organisms have been identified as bioluminescent, such as dinoflagellates, jellyfishes, copepods, euphausiids, squids and some fishes. Bioluminescence is the result of a substance as Luciferin being oxidized in the presence of a catalitic enzyme, the Luciferase. Bioluminescence is triggered by unexplained internal actions or external actions such as surface and internal waves, ship, fish and whale movements, upwelling.

Bioluminescence can be measured and collected *VISUALLY* and recorded in coded form from standard tables. Data collected includes information about Stimulus (causes such as light, wave action, rain or fish), Colour, Kind (continuous, patches, bands, blobs, shapes), Duration (seconds or continuous), and Extent (size of patches or continuous) of the Bioluminescence.

Various *MESH NETS* with preset depth catchers catch larger plankton while the smallest plankton type must be centrifugated because it is too small for the smallest mesh.



Figure 16. Shallow water plankton net

3. FURTHER READING LIST

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